

# Triangular 2D and 3D Grid Refinement for Atmosphere and Ocean Simulation

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DEKLIM Projekt Nr. 01 LD 0037

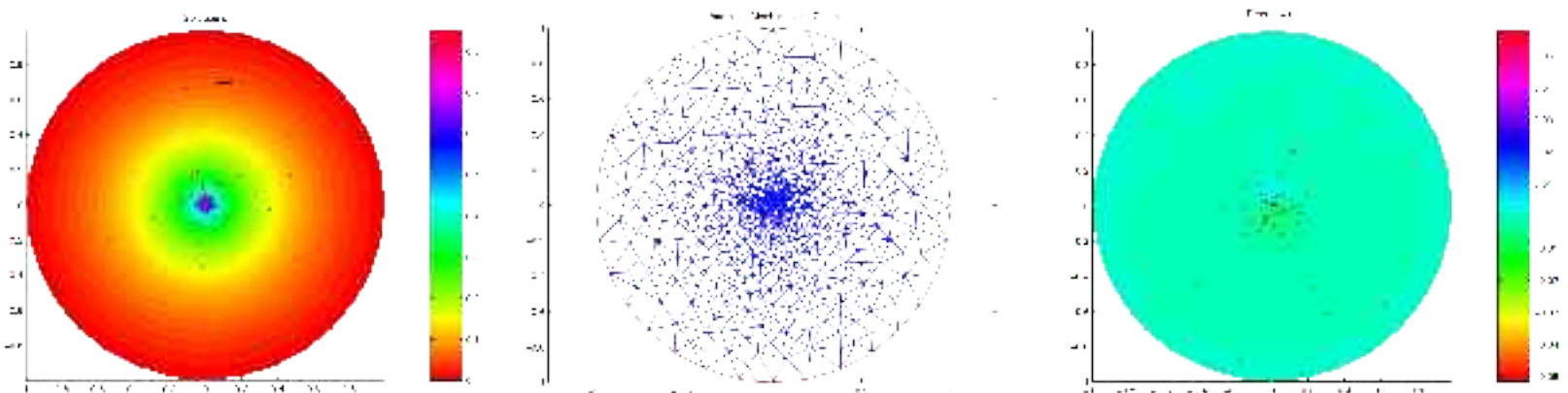


DFG-Stipendium Nr. BE2314/3-1

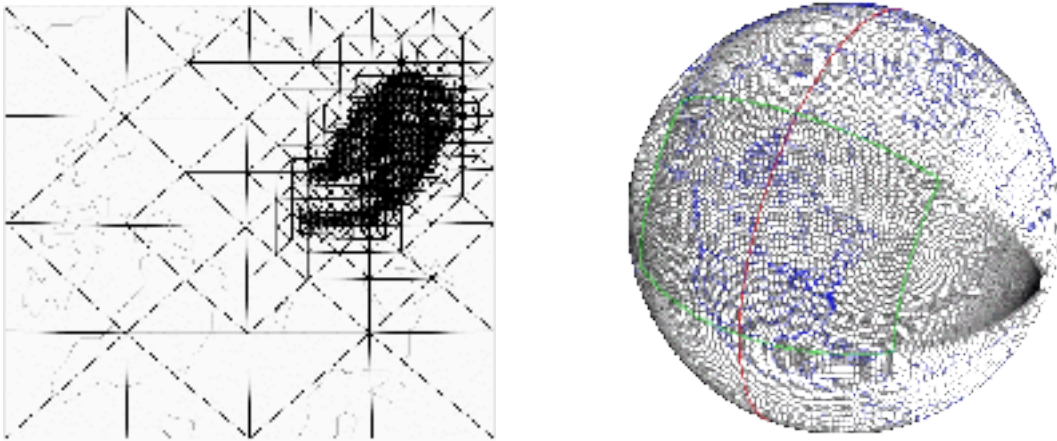


# Introduction – Perspective Adaptivity

**Equidistribution of error** → **rigorous error notation**

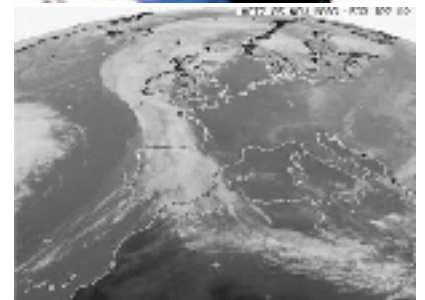
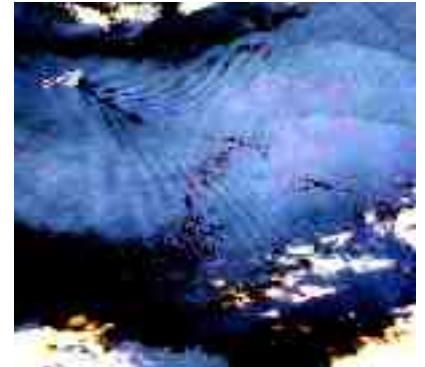


**Enhanced resolution** → **scale analysis**

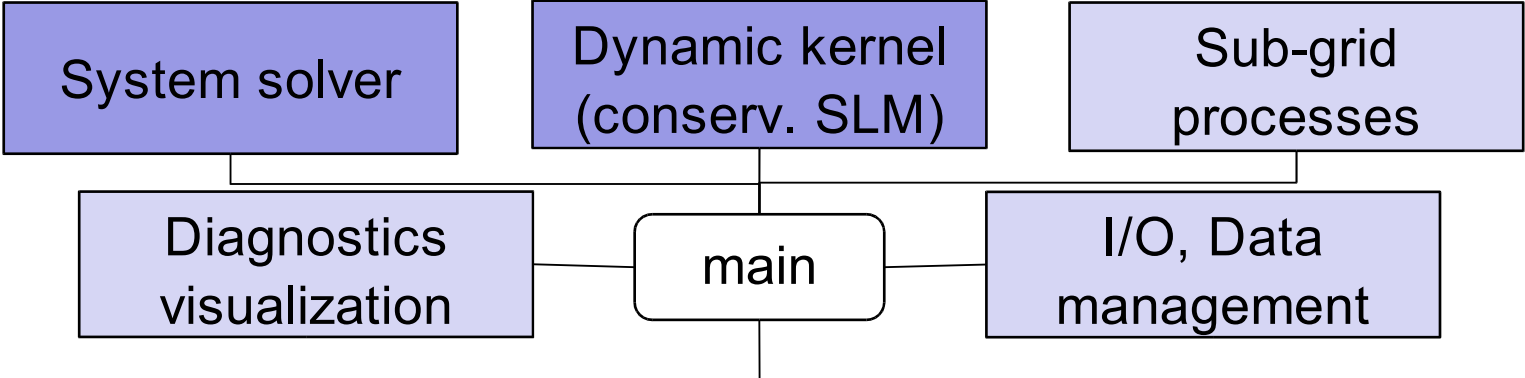
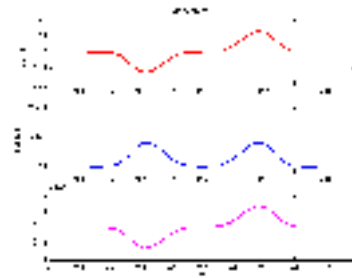
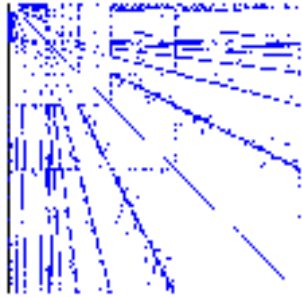


# Introduction – why adaptive modeling?

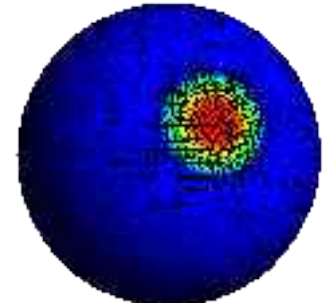
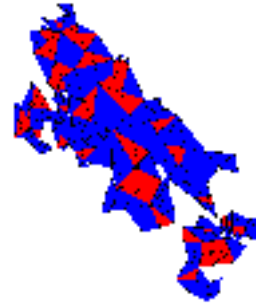
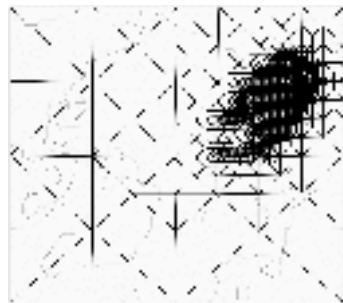
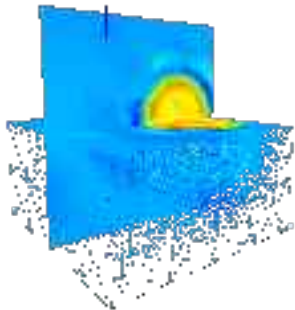
- **Scale interaction**
- **Sensitivity analysis  
(local – global scale)**
  
- **Fronts (large gradients)**
- **Embedded local phenomena**
  
- **Filamentation in tracers**
- **Point sources for tracers**
  
- **Efficient utilization of  
computing resources**



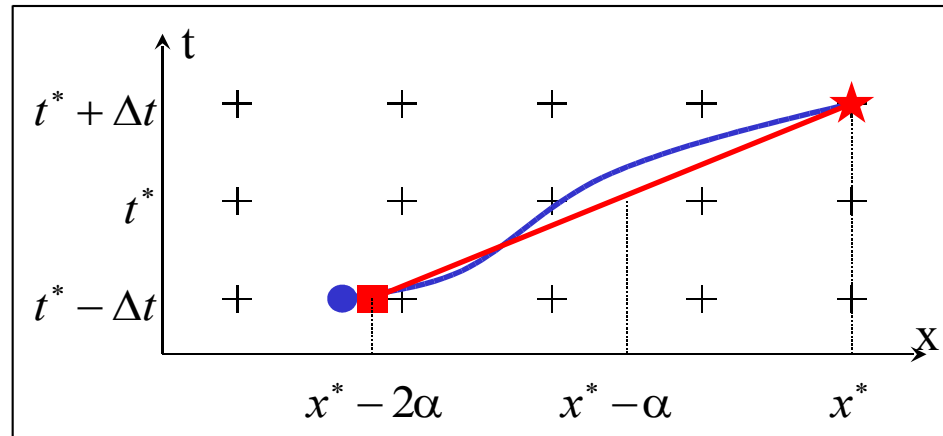
# Modular Adaptive Software



**Grid generator amatos**  
(<http://www-m3.ma.tum.de/software/amatos>)



# Discretization



$$0 = \frac{dc}{dt} \approx \frac{c(\vec{x}, t + \Delta t) - c(\vec{x} - 2\vec{\alpha}, t - \Delta t)}{2\Delta l}$$

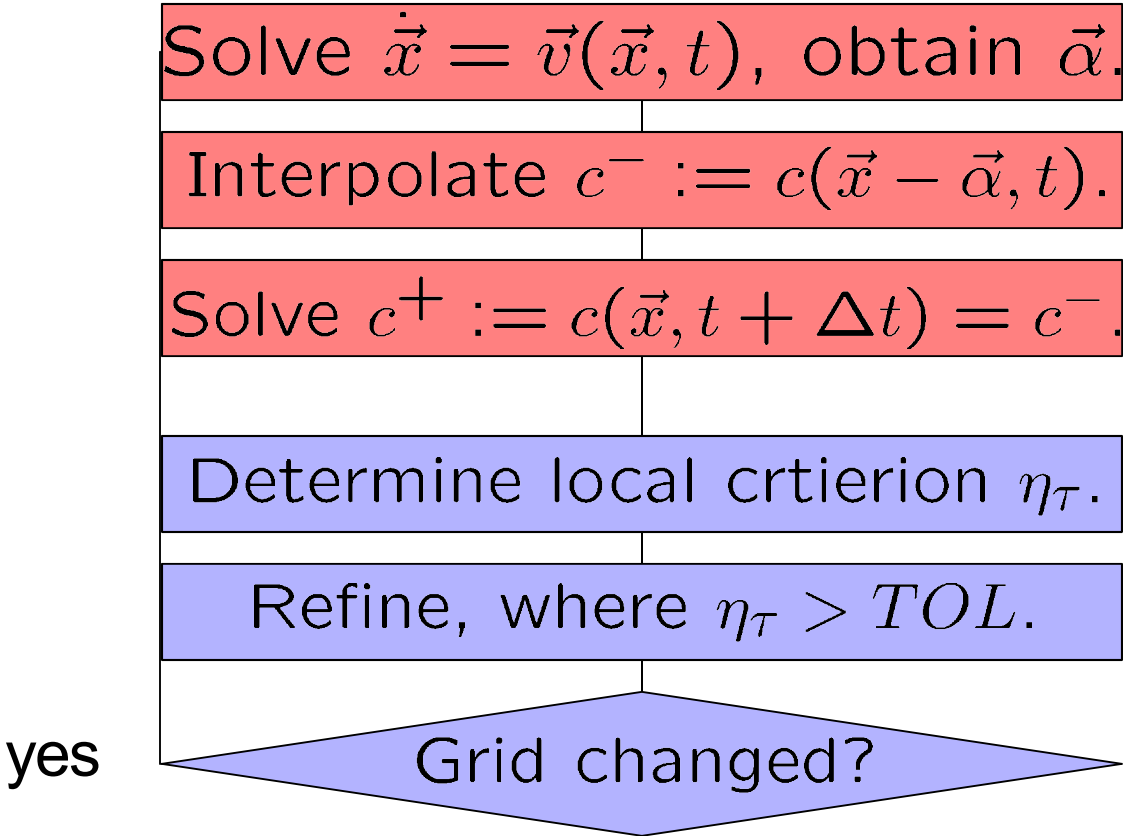
$$\Rightarrow c(\vec{x}, t + \Delta t) = c(\vec{x} - 2\vec{\alpha}, t - \Delta t)$$

$$\frac{\vec{\alpha}}{\Delta t} \approx \frac{d\vec{x}}{dt} = \vec{v}(\vec{x}, t)$$

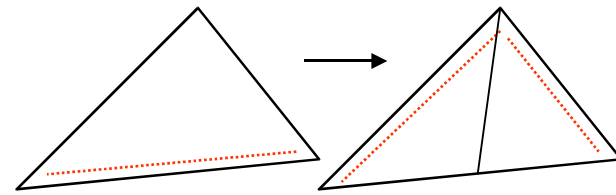
$$\text{Iteration: } \vec{\alpha}^{k+1} = dt \cdot \vec{v}(\vec{x}^* - \vec{\alpha}^k, t^*)$$



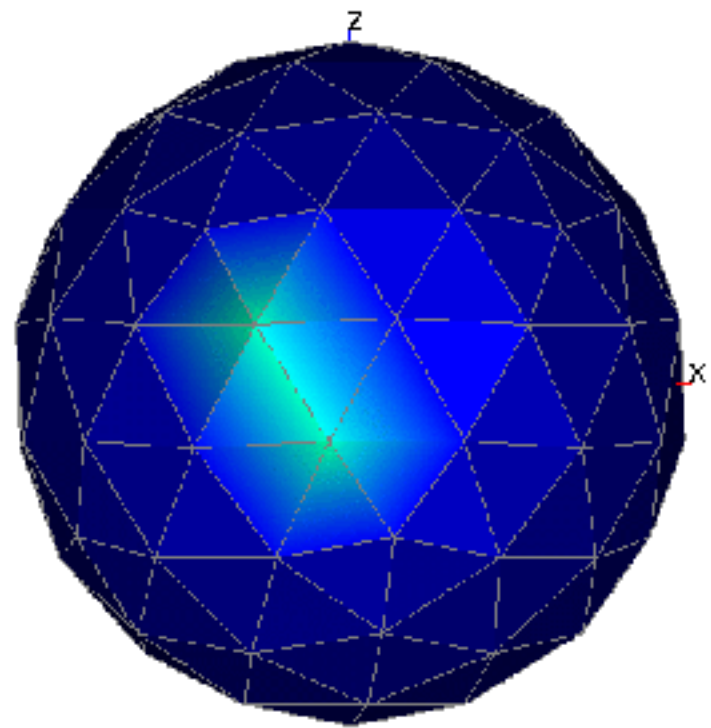
# Algorithmus: Semi-Lagrange Methode (SLM)



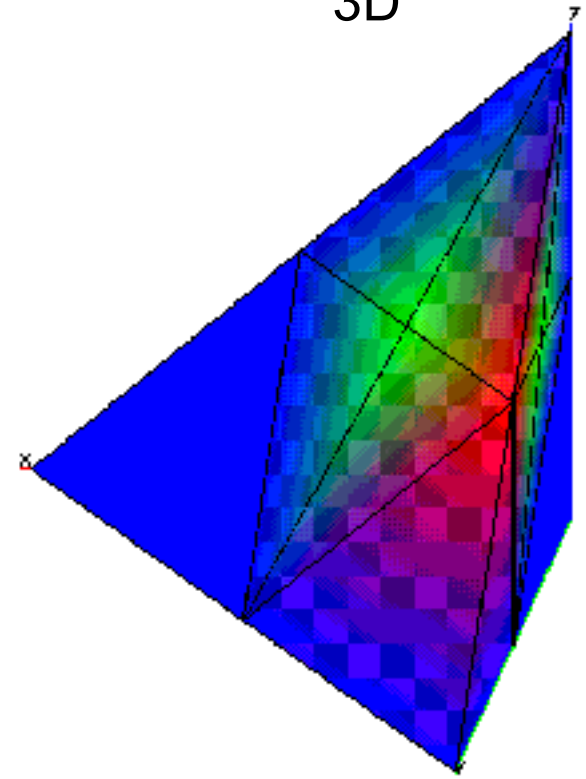
# Refinement Strategy



2D



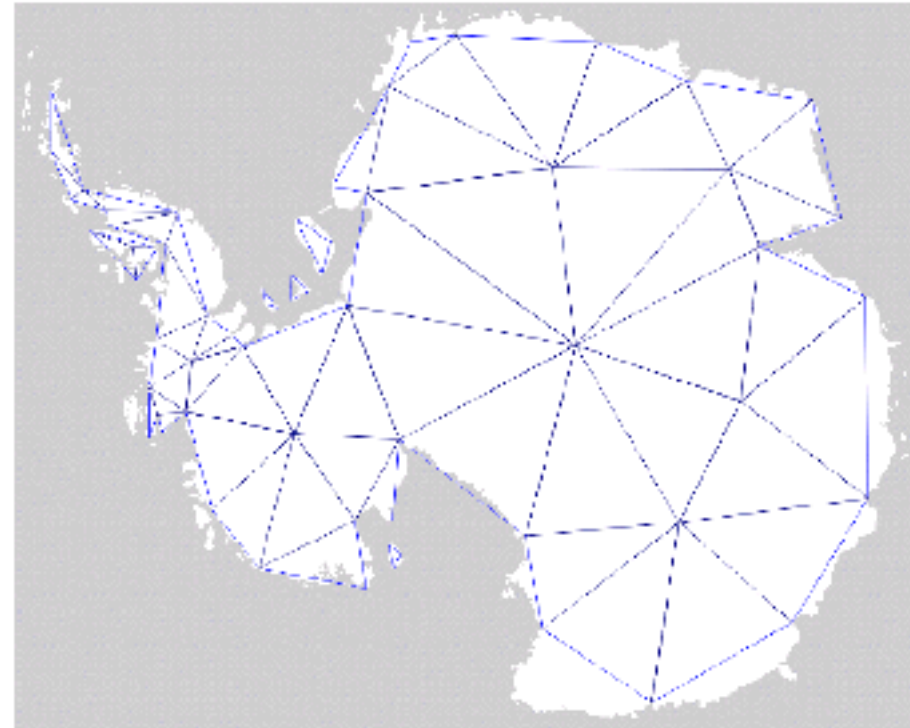
3D



# Complex Geometries

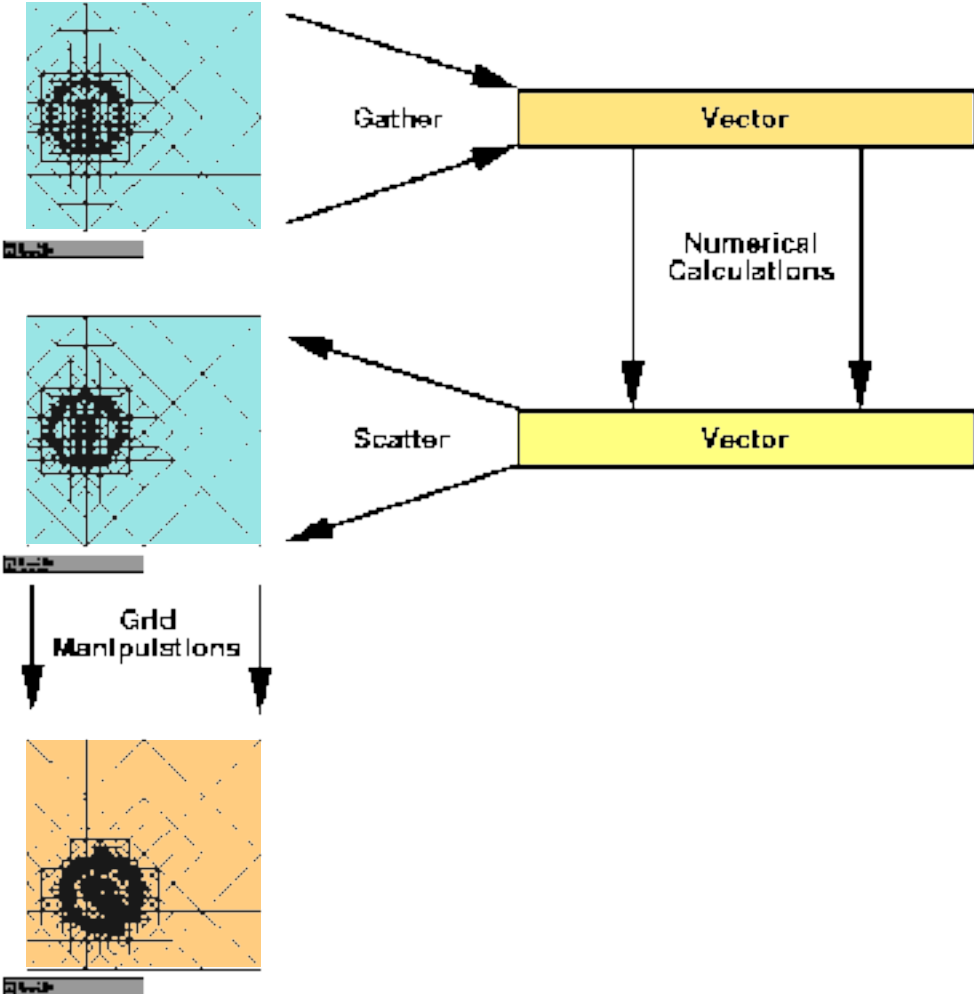


Polygonal domain

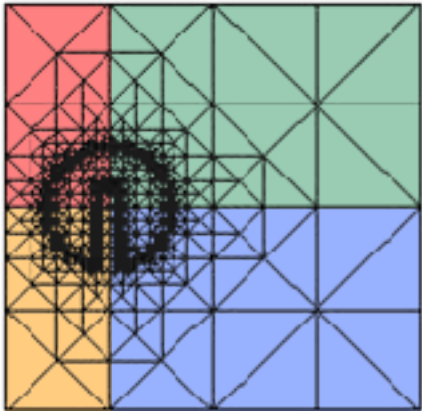


Bitmapped domain

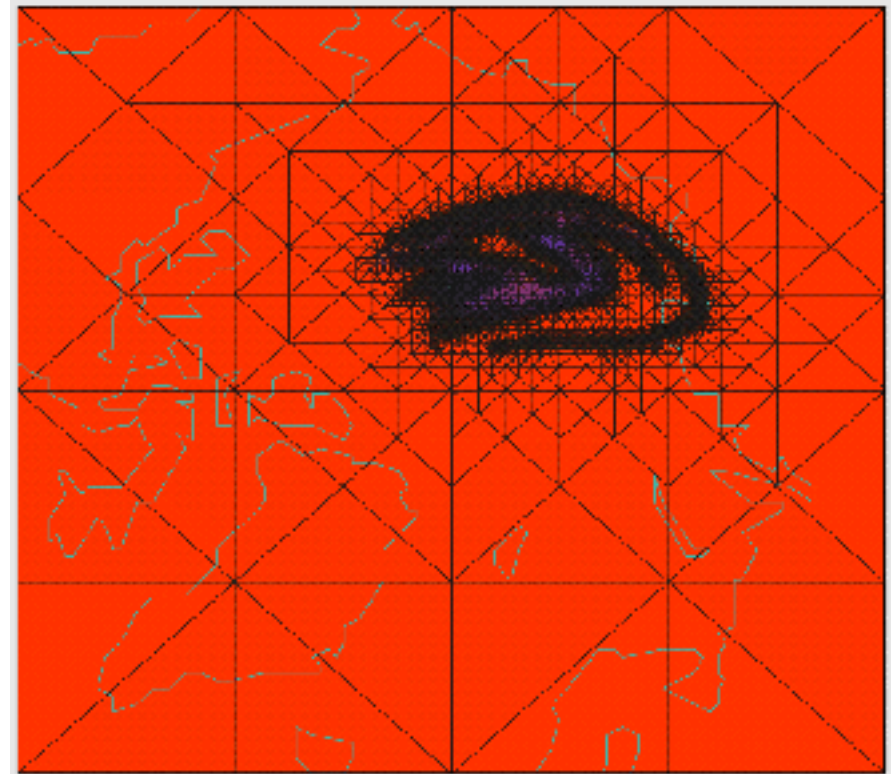
# Data Management and Numerics



# Data Management and Parallelization



# Parallelization



## Partitioning problem

- Distribute cells in equally sized sets (partitions)
- Partitions shall be connected
- Partitions have to be re-calculated frequently
- Data movement has to be minimized
- Algorithm has to be parallel/low computational effort



# Parallelization

## Excursion: Space-filling curves

**Question:**  $\exists$  surj.  $\wedge$  cont. map  $\mathbb{R} \rightarrow \mathbb{C}$ ?

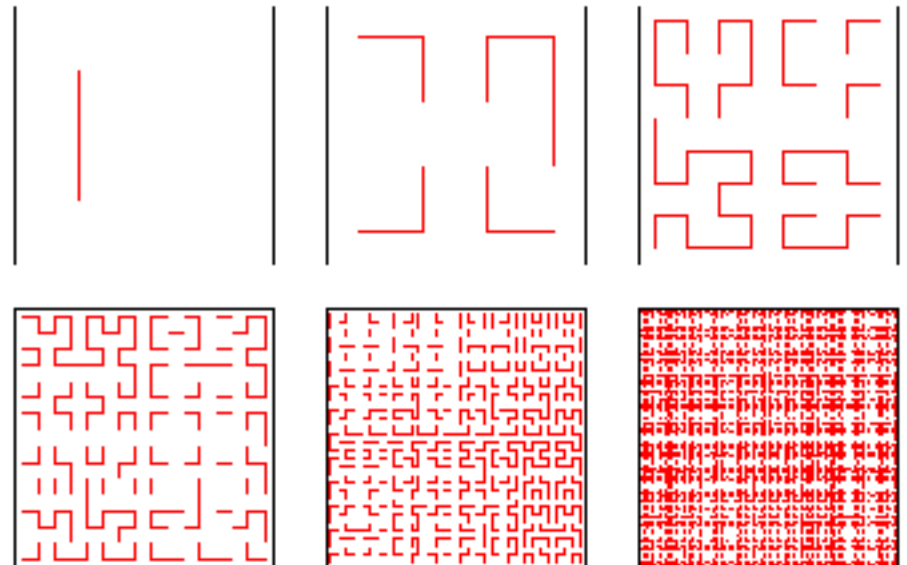
**Answer:** Yes! Constructive recursive proof.



Giuseppe Peano  
(1858-1932)

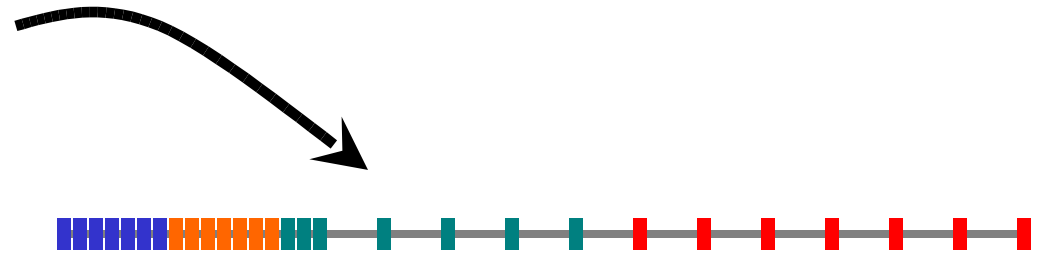
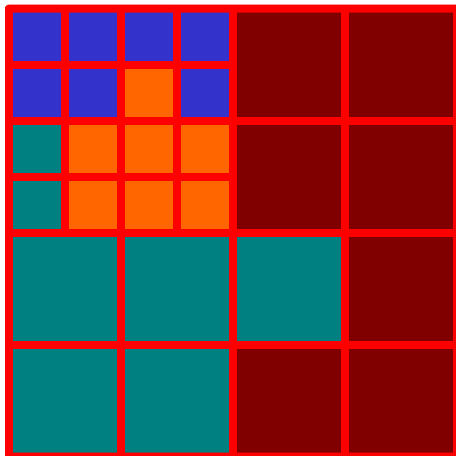
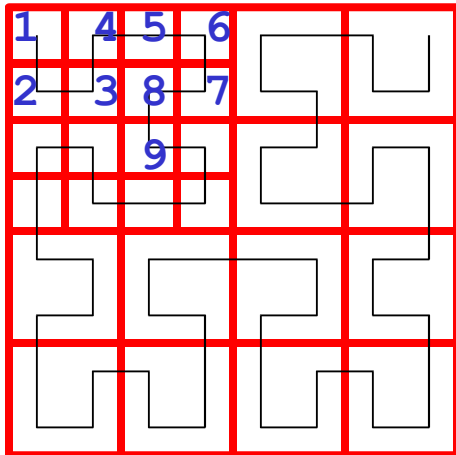


David Hilbert  
(1862-1943)



# Parallelization

## Space-filling curve for load balancing

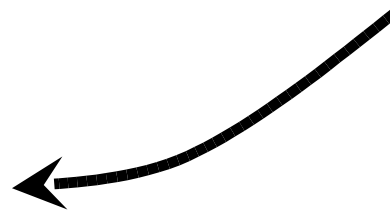


Proc.  
1

Proc.  
2

Proc.  
3

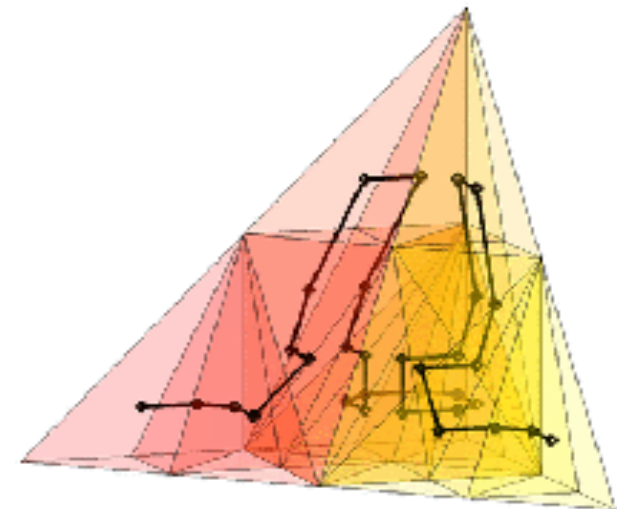
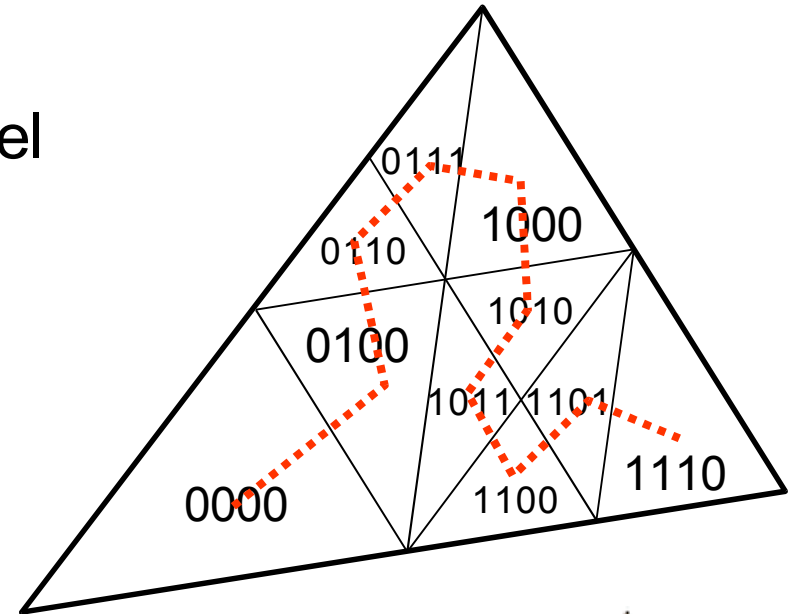
Proc.  
4



# Parallelization

## Algorithm for triangles

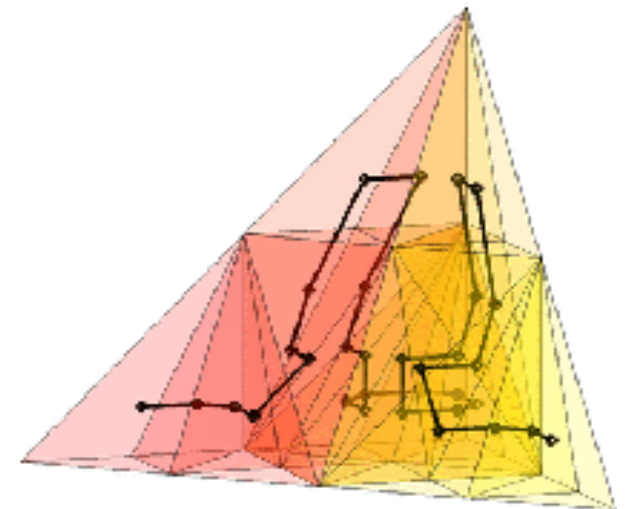
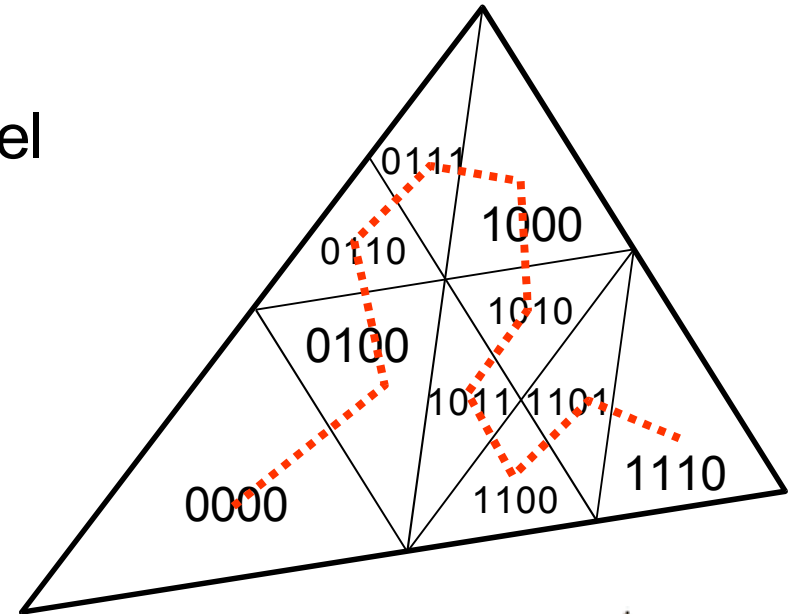
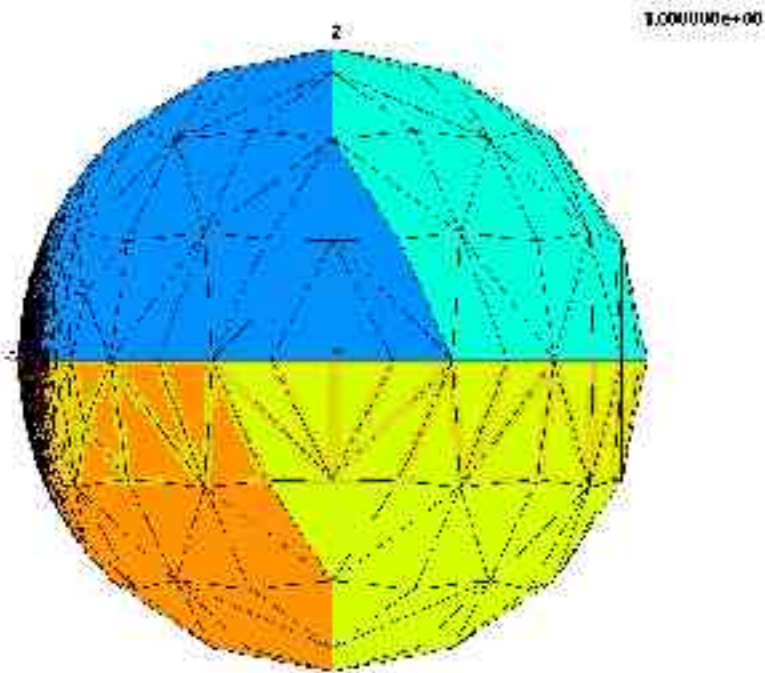
1. One bit per refinement level
2. Set bits while refining



# Parallelization

## Algorithm for triangles

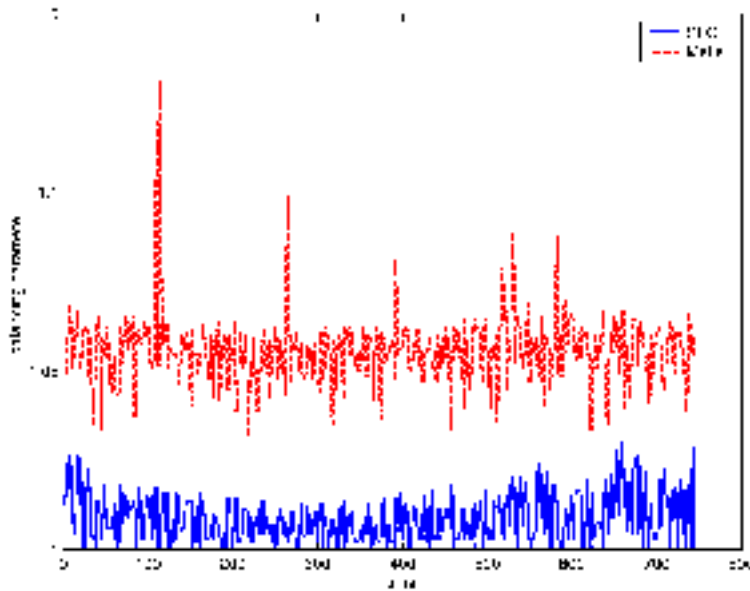
1. One bit per refinement level
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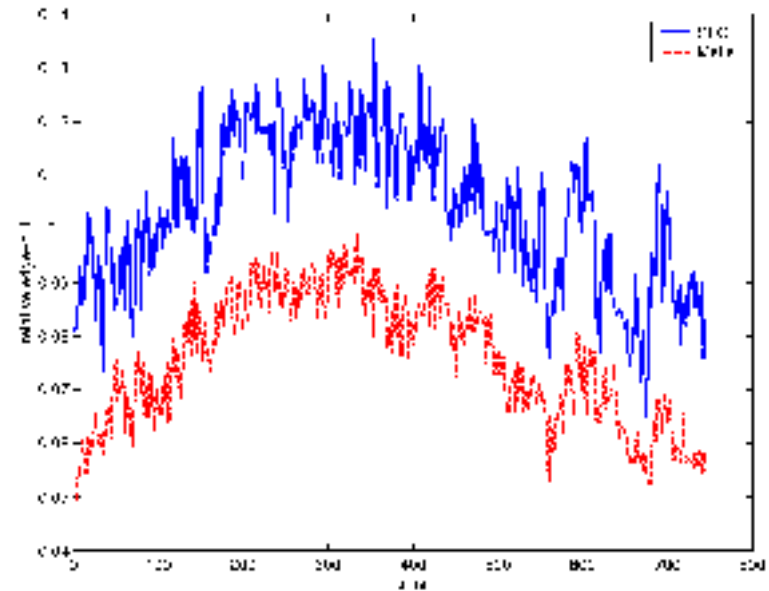
# Results: Tracer Advection

## Artificial tracer in Arctic stratosphere

Load balancing

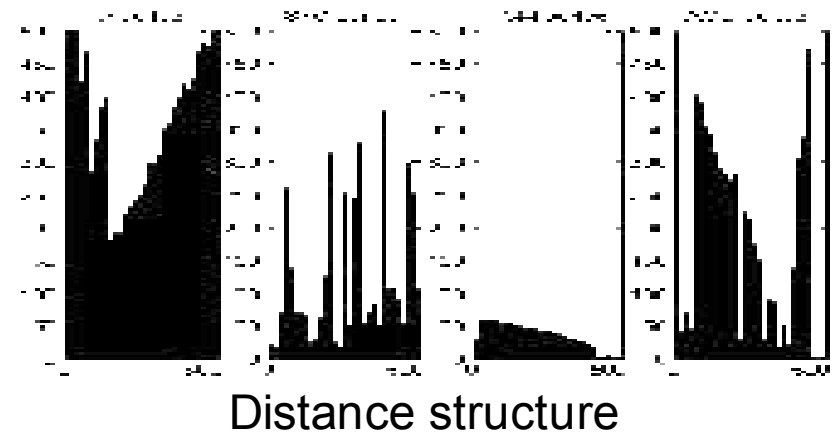
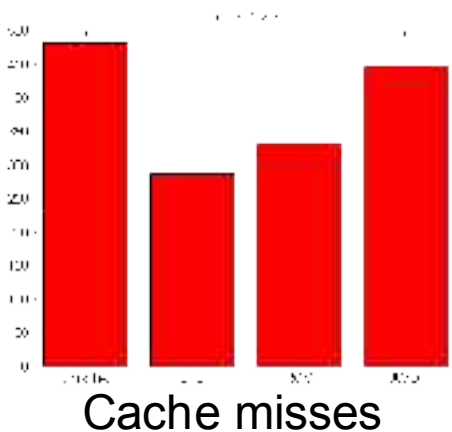
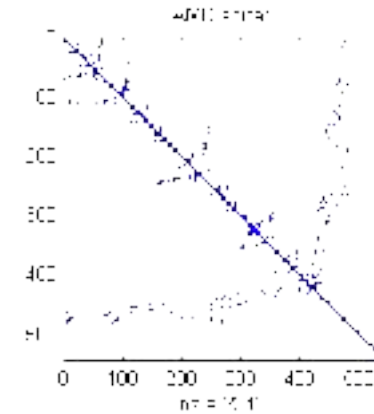
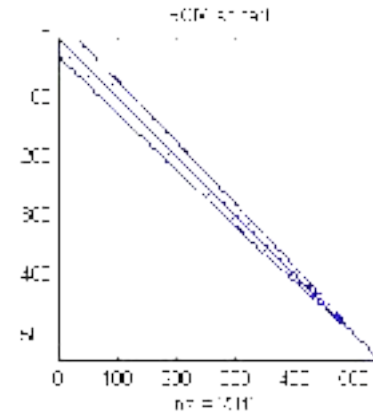
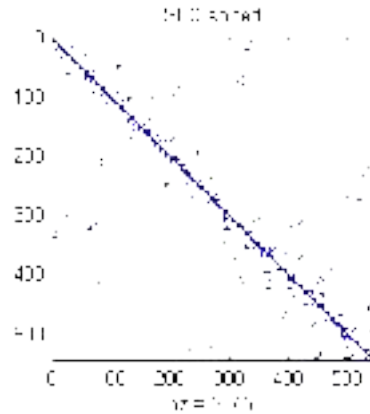
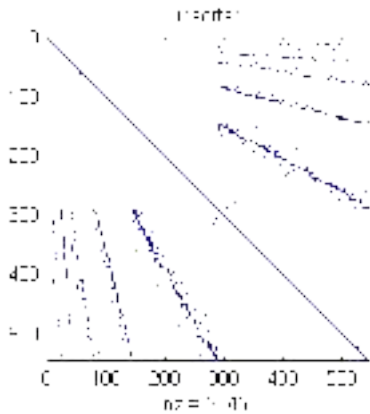


Edge-cut



# Data Management revisited

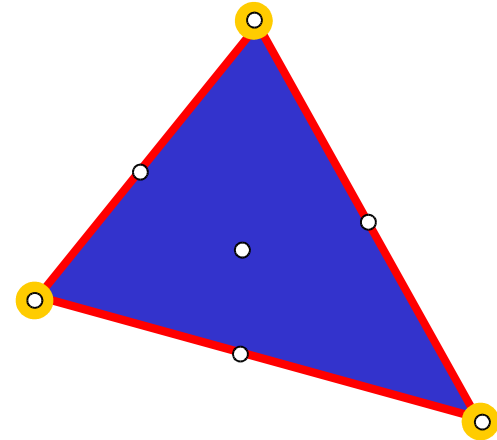
## Connectivity matrix with different orderings



# FEM support

Main data objects:

**nodes**, **edges**, **triangles**



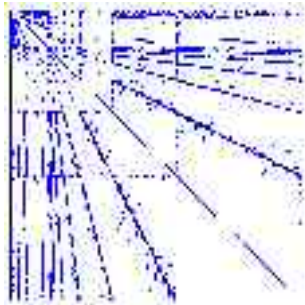
FEM-Signature:

- Unknowns on **nodes**
- Unknowns on **edges**
- Unknowns on **triangles**
- Position in barycentric coordinates  
(for **edges** and **triangles**)

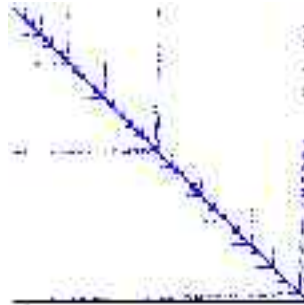


# Space-Filling Curves: Matrix Ordering

## Structure of matrix



tree-sorted



quotient  
minimum degree

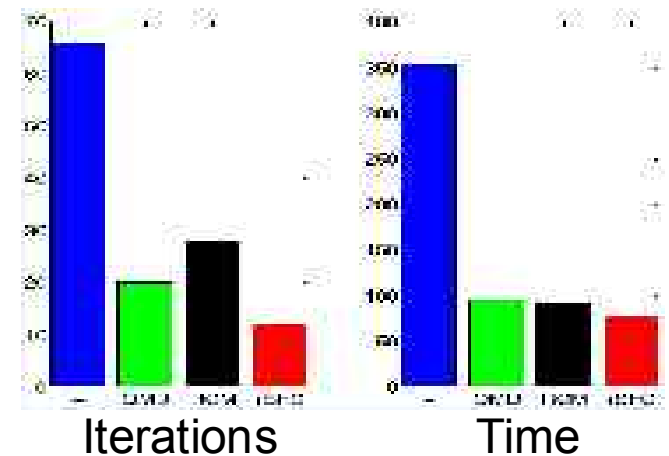


reverse  
Cuthill-McKee



reverse SFC

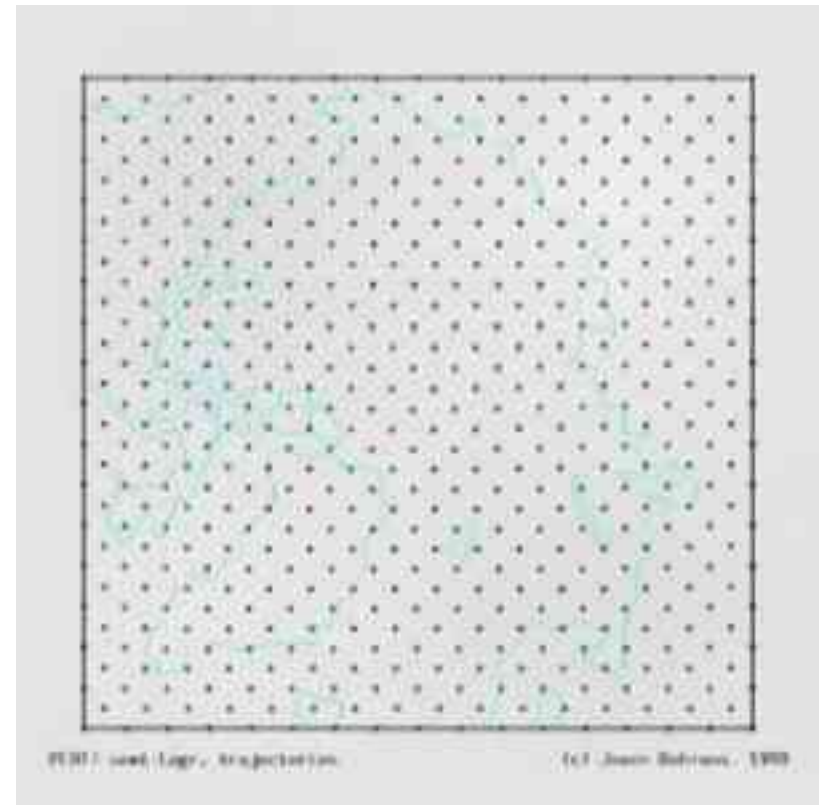
- System with ~200.000 unknowns
- Utilizes preconditioned BiCGStab
- ILU pre-conditioning



# Example 1: linear advection

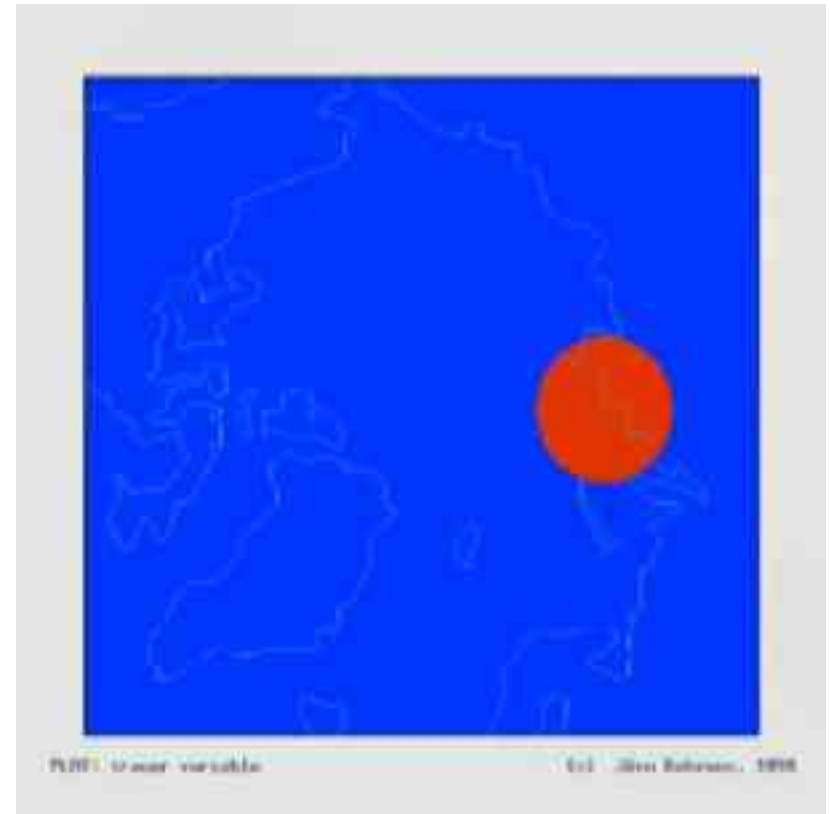
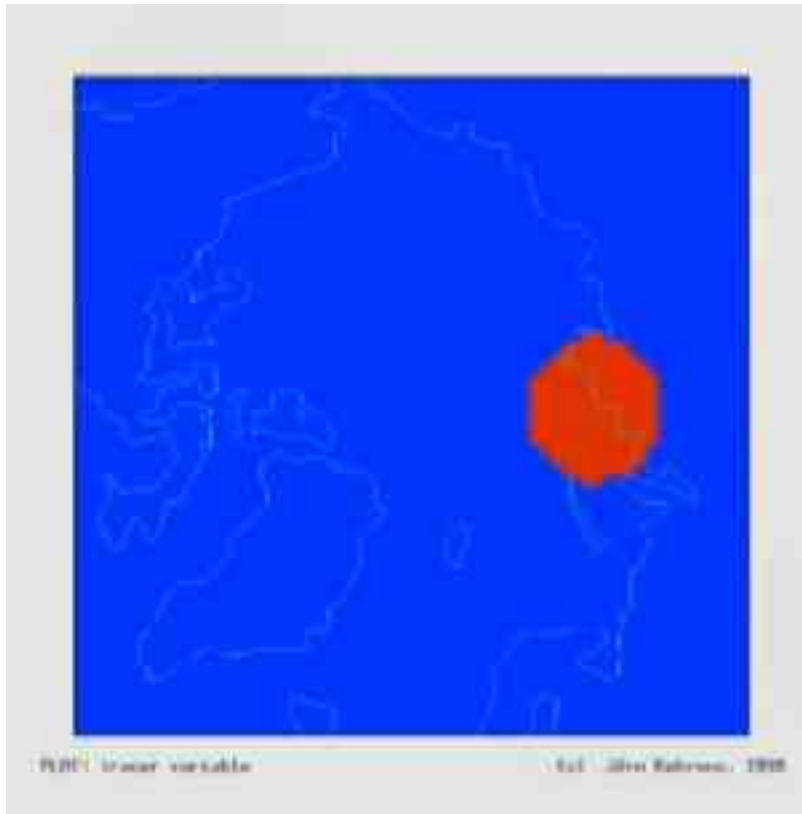
## Simulation of tracer transport

Resolution of wind data:  
50 x 50 km  
Situation in January 1990,  
70 hPA layer (18.000 m)



# Example 1: linear advection

## Simulation of tracer transport



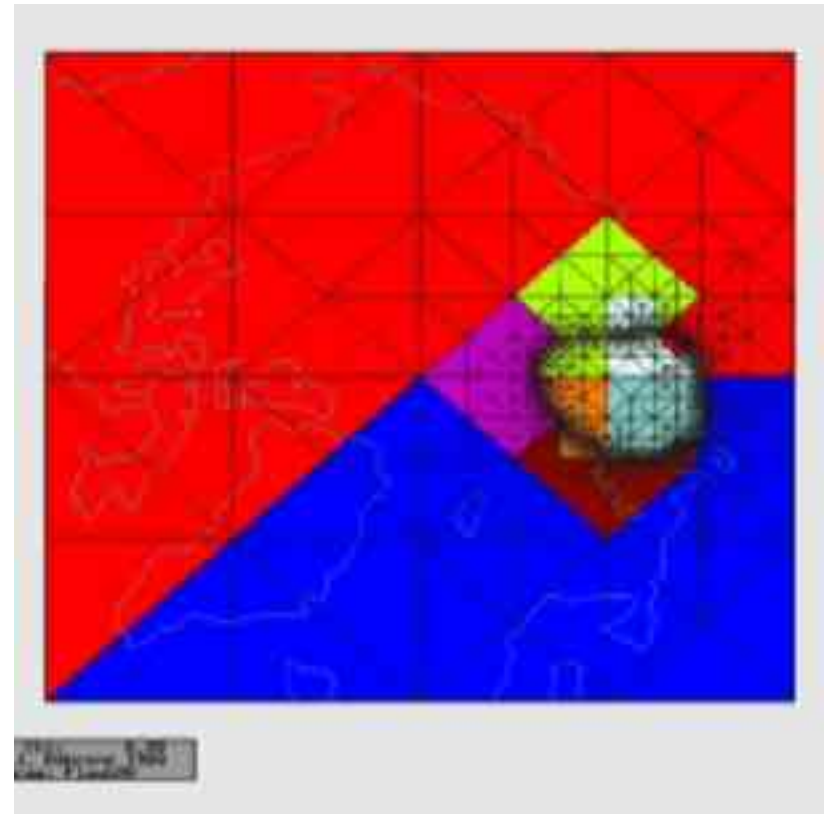
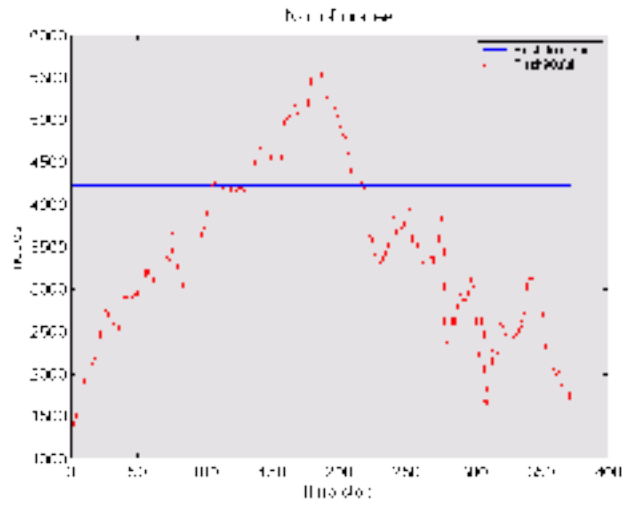
Resolution: 50 km uniform

Resolution: 5 km local

# Example 1: linear advection

## Simulation of tracer transport

Costs:  
Uniform vs. adaptive



# Example 2: shallow water equations

## Flow over isolated mountain

$$\begin{aligned} \frac{d\zeta}{dt} + \zeta\delta + f\delta &= \bar{u} \cdot \nabla_S f \\ \frac{d\delta}{dt} + \Delta_S \Phi - f\zeta &= (\mathbf{n} \times \mathbf{u}) \cdot \nabla_S f - J(\mathbf{u}) \\ \frac{d\Phi}{dt} &= (\Phi - \Phi_0)\delta. \end{aligned}$$

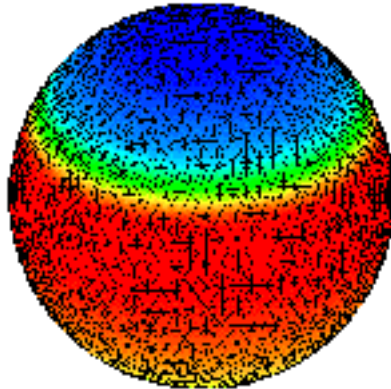
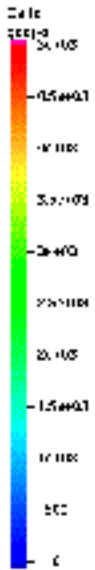
Equations in vorticity-divergence form

$\zeta := \text{rot } \vec{V}$  (vorticity)

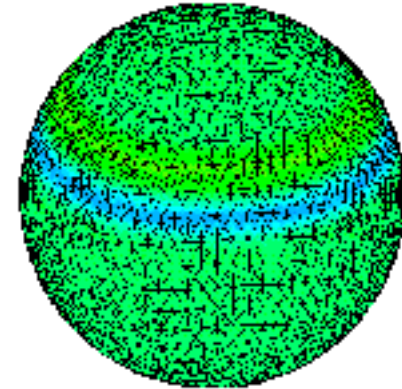
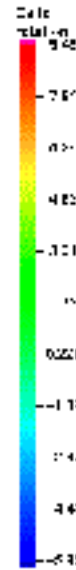
$\delta := \text{div } \vec{V}$  (divergence)

0.010106H-011

0.010106H-011



Geopotential



Vorticity



# Example 3: Inverse Modeling

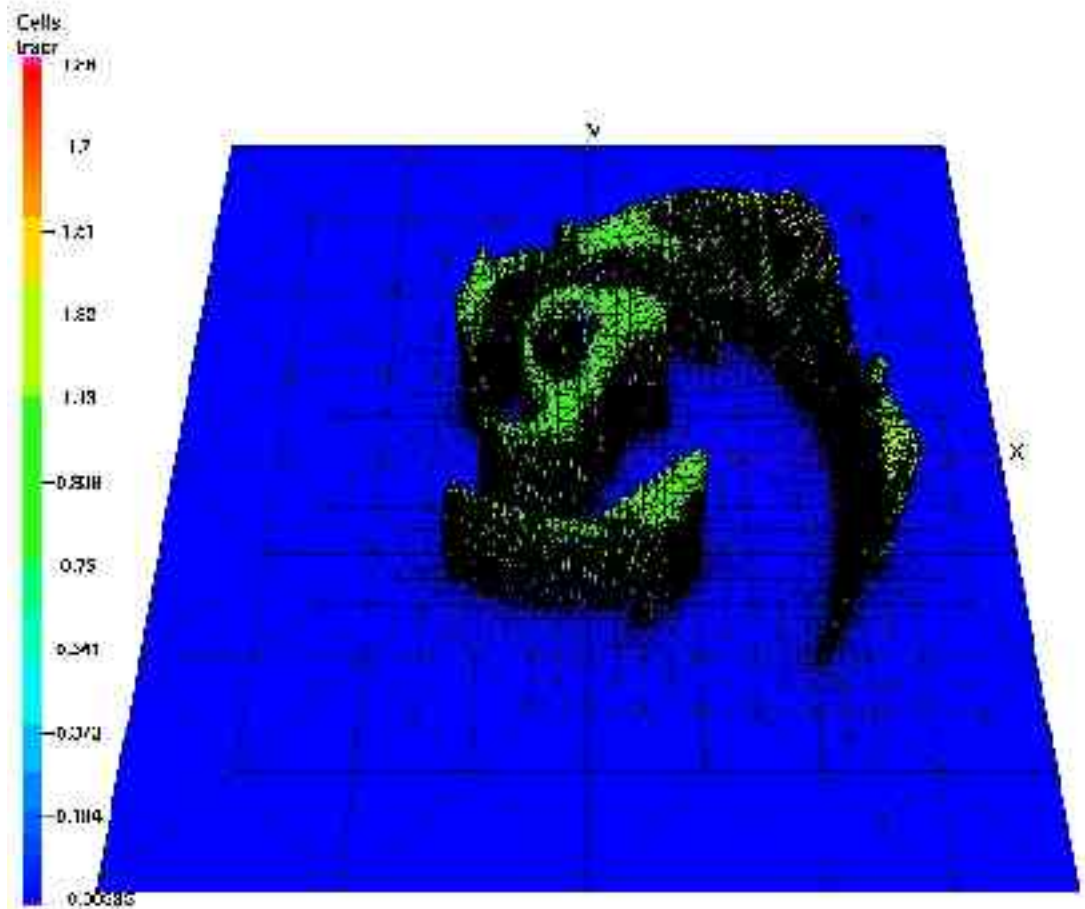
## Problem:

### Given:

- wind
- tracer density distribution

### Question:

source  
of tracer?



# Conclusions

- Triangular grid generation for simplicity and complex domains
- Adaptive grid refinement for accuracy and efficiency
- SFC for partitioning in parallel applications
- SFC ordering for efficient data access and matrix reordering
- Examples from tracer transport to dynamical core

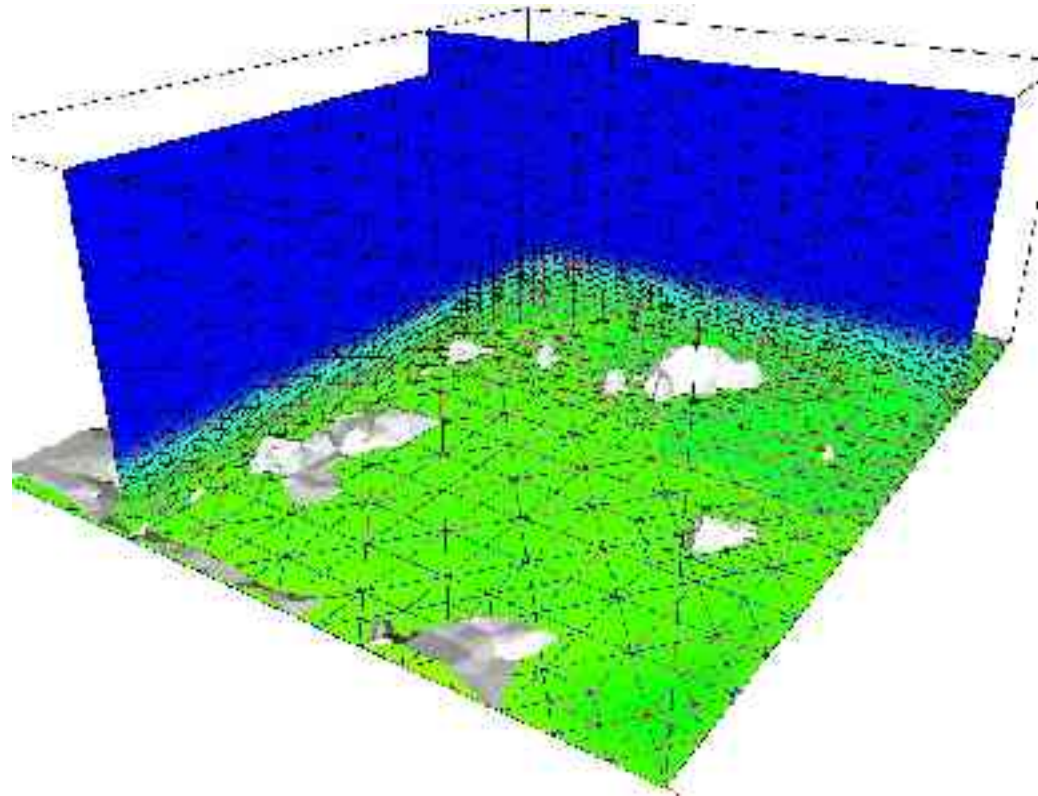


## Atmosphere/Ocean is 3D

### Tracer transport in 3D

Wind Data from  
measurements

CO density from  
measurements (in ppb)



### Future topics:

- Error estimation/ refinement criteria
- More realistic problems
- Coupling, parameterization, ...

